# Biostratigraphic Aspects and Depositional Environments of Eocene Deposits of Northwestern Suluova (Amasya, Northern Turkey)

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## Abstract

The Eocene deposits which are interpreted as deposits of the northern branch of Neo-Tethys crop out over a wide area in northern Turkey, in an approximately E-W belt. The biostratigraphic characteristics and depositional environments of the Celtek and Armutlu formations in Suluova (Amasya) were investigated. Four stratigraphic sections were measured and analyzed. Eight planktic foraminiferal genera with fourteen species and thirty benthic foraminiferal genera with five species were identified. Two planktic foraminiferal biozones from the Early Eocene (Morozovella aragonensis and Acarinina pentacamerata biozones) and one benthic foraminiferal biozone from the Middle Eocene (Nummulites laevigatus biozone) were defined. The charophytes (with Nitellopsis and Harrisichara) associated with these sediments can be probably attributed to the Disermas-Piveteaui zones. These biozones establish the age of the Armutlu Formation as Early-Middle Eocene. Sedimentological and paleontological data indicate that the Celtek and Armutlu formations were deposited in delta, lagoon, and shallow marine environments.

Keywords: Biostratigraphy, Depositional environments, Eocene, Foraminifera, Charophytes, Amasya-Turkey.

# 1. Introduction

The Eocene deposits crop out over a wide area of the Cankırı–Corum Basin of northern Turkey, in an approximately E-W belt. The Armutlu Formation is interpreted as deposits of the north branch of Neo-Tethys Figs. 1, 2). Brown coals located in the study area are quality within brown coals Turkey has. They have an important role for area economy. Therefore, the areawas measured across the exposure belt along with detailed paleontological sampling for age determinations.



Figure 1. a-Location map of the study area, b-Position of the Armutlu and Celtek Formation within Neo-Tethys realm (Sengor and Yilmaz, 1983)



# 2. Material and methods

Four measured sections were analyzed in the study area. The systematic samples were collected from the measured section the samples which were collected through the measured sections range from 50 cm to 15 m. From these samples, benthic and planktic foraminifera and charophytes were identified. Quantitative analyses of benthic foraminifera were carried out on 54 samples in total of which 23 were rock thin sections and 31 were soft rock samples. In addition some of the selected soft rock samples were prepared oriente (18 individual) sections. Soft rock samples were washed through a 63  $\mu$ m sieve using a 17% hydrogen peroxide solution for 24 hours. The remaining fraction was oven-dried and dry sieved at 63  $\mu$ m, 125  $\mu$ m and 250  $\mu$ m. In the end of micropaleontological examinations, 19 species and 38 genera of foraminifera were identified at all localities. Some representatives of every investigated species were examined with a scanning electron microscope (SEM).

### 3. Biostratigraphy

The Celtek and Armutlu formations are generally dated as Early-Middle Eocene. A more precise age is determined through the measurement of four sections in the study area (Figs. 3-7). Two planktic and one benthic foraminiferal biozones were identified in the calcite-cemented mudstone and sandstone as well as in the marl alternations of the Armutlu Formation. Determination of the foraminiferal genera and species was based on Ellis and Messina, 1966; Beckmann et al., 1969; Bolli, 1957a, b, 1966; Ejel, 1967; Postuma, 1971; Toumarkine, 1978 and Toumarkine and Luterbacher, 1985; Serra-Kiel et al., 1999.

#### Foraminifera

Planktic foraminifera include: Globorotalia sp., Planorotalites sp., Acarinina pentacamerata (SUBBOTINA), Acarinina primitiva (FINLAY), Acarinina soldadoensis soldadoensis (BRONNIMANN), Morozovella acuta (TOULMIN), Morozovella aequa (CUSHMAN and RENZ), Morozovella aragonensis (NUTTALL), Morozovella caucasica (GLAESSNER), Morozovella formosa formosa (BOLLI), Morozovella subbotinae (MOROZOVA), Truncorotaloides sp., Pseudohastigerina wilcoxensis (CUSHMAN and PONTON), "Globigerinoides" higginsi BOLLI, Globigerina inaequispira SUBBOTINA, Globigerina linaperta FINLAY, Globigerina velascoensis CUSHMAN, Hastigerina sp..



Figure 3. Sirikli Hill measured section

Benthic foraminifera include: Ammodiscus sp., Vulvulina sp., Pseudobolivina sp., Verneuilina sp., Dorothia sp., Semivulvulina sp., Textularia sp., Austrotrillina sp., Dentalina sp., Nodosaria sp., Frondicularia sp., Cribrolenticulina sp., Lenticulina sp., Robulus sp., Marginulinopsis sp., Percultazonaria sp., Hemirobulimina sp., Marginulina sp., Epistomina sp., Bolivina sp., Praebulimina sp., Bulimina sp., Buliminoides sp., Planorbulina sp., Anomalina sp., Anomalinoides sp., Daviesina sp., Nummulites gallensis (HEIM), Nummulites laevigatus (BRUGUIERE), Nummulites lehneri SCHAUB, Nummulites uranensis (de la HARPE), Nummulites sp., Operculina subgranulosa d'ORBIGNY, Discocyclina sp..

In addition fossil ostracodes and gastropods, such as Bayania sp. and Burtinella sp. were also found.



Figure 4. New Celtek measured section

#### Charophytes

Three samples of the Tersakan member yielded charophyte gyrogonites. The sample A0.1 contains several small gyrogonites of *Peckichara vel Sphaero*chara.

*Peckichara vel Sphaero*chara (sample A0.1) spheroid gyrogonites are 540 x 600 mu in length; 540 x 600 mu in size. The ISI is 0.98-1.03. A 7-8 plano-convex convolution and very weak ornamentation (some disseminated points) were noted. The cells are not thinner and narrower on the apex, the apical nodules are absent. Basal pore is 40 to 60 mu in diameter. It is surrounded by convex nodules. The basal plate is thick and straight (80 mu diam. and 60 mu thick). These forms could be attributed to *Peckichara* or *Sphaerochara*. The size and basal plate plaid in favor of *Sphaerochara* (as for example *Sph. edda* or *Sph. hirmeri*) but the general shape and cell ornamentation is very similar to *Peckichara disermas* or *P. piveteaui*.

In samples A0.2 and A0.5, there are several big forms clearly attributed to *Nitellopsis* sp. It is an ovoid gyrogonite of 1080-1180 mu in length and 920-960 mu in size. The ISI is 1.12-1.28; there are 8-9 plano-convex convolutions; the cells are only slightly thinner and narrower on the apex, the apical nodules are prominent. The base forms a small column. This taxon is a *Nitellopsis*; generally the older forms of this genus (Paleocene and



Figure 5. Kalayli Hill measured section

Eocene) are ornamented (*N. thaleri*, *N. dutemplei*) or very big (*N. major*, *C. helicteres*). The number of gyrogonites is too small to give a specific species level here, but they clearly represent small and unornamented forms of *Nitellopsis*. They show some affinities with N.C. sigali (ornamentation and size), but the number of convolutions is higher in sigali (9-11) than in the Celtek species (6-8).

In sample A02, one badly preserved *Harrisichara* is identified. The small size and the ornamentation indicate perhaps *Harrisichara triquetra*, but it could be also a small example of *H. tubercu*lata.

#### **3.1.** The Benthic Foraminiferal Biozone

#### 3.1.1. Zone I, *Nummulites laevigatus* Range Zone (=SBZ13)

Age: Middle Eocene (Early Lutetian)

The lower boundary of this biozone has been defined by the first occurrence of *Nummulites laevigatus* (BRUGUIERE). The upper boundary has been determined by the disappearance of *Nummulites laevigatus* (BRUGUIERE). In addition to the markers of this zone, several foraminiferal species have been recorded: *Nummulites gallensis* (HEIM), *Nummulites lehneri* SCHAUB, *Nummulites uranensis* (de la HARPE).

The Nummulites laevigatus Range Zone has been recognized in the Sirikli Hill section in the study (Fig. 3).

#### **3.2. The Planktic Foraminiferal Biozones**

**3.2.1.Zone I**, *Morozovella aragonensis***Zone** (=P8 zone of TOUMARKINE and LUTERBACHER, 1985, P6b sub-zone of BERGGREN et al., 1995 and P7 zone of OLSSON et al., 1999)

Age: Early Eocene

Author: BOLLI (1957a, b)

This biozone is defined here as the partial range from the last occurrence datum (LOD) of *Morozovella* formosaformosa (BOLLI) to the first occurrence datum (FOD) of Acarinina pentacamerata (SUBBOTINA).

This zone is defined as the interval between the first occurrence of *Acarinina pentacamerata* (SUBBOTINA) and the last occurrence of *Morozovella formosa formosa* (BOLLI). In addition to the markers of this zone, several foraminiferal species have been recorded: *Ammodiscus* spp., *Verneuilina* sp., *Vulvulina* sp., *Semivulvulina* sp., *Textularia* sp., *Nodosaria* spp., *Dorothia* spp., *Lenticulina* spp., *Robulus* spp., *Acarinina pentacamerata* (SUBBOTINA), *Acarinina primitiva* FINLAY, *Acarinina soldadoensis* soldadoensis(BRONNIMANN), *Morozovella aragonensis* (NUTTALL), *Morozovella caucasica* (GLAESNER), *Morozovella formosa formosa* (BOLLI), *Globigerina inaequispira* SUBBOTINA, *Globigerina linaperta* FINLAY, *Praebulimina* sp. *Acarinina pentacamerata* 20ne conformably overlies this zone. It coincides with the P8 zone of TOUMARKINE and LUTERBACHER, 1985; P6b sub-zone of BERGGREN et al., 1995 and P7 zone of OLSSON et al., 1999. It is also equivalent to that identified by BECKMANN et al., 1969; POSTUMA, 1971. According to the abovementioned records an Early Eocene age is suggested to this biozone.

The Morozovella aragonensis Zone has been recognized in the New Celtek section, in this study (Fig. 4).

**3.2.2.** Zone II, *Acarinina pentacamerata* Zone (=P9 zone of TOUMARKINE and LUTERBACHER, 1985)

Age: Early Eocene

Author: KRASHENINNIKOV (1965) as Subzone

This zone is defined as the first occurrence of *Acarinina pentacamerata* (SUBBOTINA) and "Globigerinoides" higginsi BOLLI and the disappearance of Morozovella aragonensis (NUTTALL) at the bottom of the zone. In addition to the markers of this zone, several foraminiferal species have been recorded: Vulvulina sp., Semivulvulina spp., Textularia sp., Nodosaria spp., Dorothia spp., Lenticulina spp., Robulus spp., Epistomina sp., Planorotalites sp., Acarinina pentacamerata (SUBBOTINA), Acarinina primitiva FINLAY, Acarinina sp., Morozovella sp., Globigerina inaequispira SUBBOTINA, Globigerina linaperta FINLAY, Globigerina sp., "Globigerinoides" higginsi BOLLI, Hastigerina sp., Bolivina sp., Bulimina sp., Siphoninella sp., Anomalina sp., Anomalinoides sp.. This zone coincides with the P9 zone proposed by TOUMARKINE and LUTERBACHER, 1985. It is also equivalent to that defined by KRASHENINNIKOV, 1965.

The Acarinina pentacamerata Zone has been recognized in the New Celtek section in this study.



Floure 6. Armutiu Village measured section



Figure 7. Biostratigraphic correlation of Celtek and Armutlu formations (modified from Koc 2002). For location of sections see Figure 3.

### 4. Sedimentology and Paleoecology

The Celtek Formation consists of delta plain and delta front facies associations. The delta plain deposits include organic rich grey mudstone and coal while delta front deposits have compositions of benthic foraminiferous greygreen mudstones, siltstone, stratified sandstone and lenticular body conglomerate. The coals were deposited in pond water and the delta swamp plain (Fig. 5).

The Armutlu Formation consists of lagoon, shoreface and offshore facies associations (Fig. 6). The lagoon facies consists of sandstone, red mudstones containing benthic foraminifera, and grey-green mudstone containing charophytes and coal. The coal was formed in a lagoon marsh. The Shoreface facies is characterized by large scale cross-bedded quartz arenitic sandstones. These sandstones commonly include bioturbation and occasionally include benthic foraminifera. The Offshore-transition facies association is composed of an alternation of parallel laminated sandstones, waved-bedded sandstones, grey-green mudstones and red mudstone facies. The Offshore facies associations consist of grey-green mudstones which include fine grained sandstone levels. There is lateral-vertical transition between the Celtek and Armutlu formations (Koc, 2002; Koc and Turkmen, 2002; Koc et al., 2002) (Fig. 8).

Gastropoda such as *Bayania* sp. and *Burtinella* sp. indicate a littoral environment. Dominance of benthic foraminifera indicates a very shallow marine paleoenvironment. The *Charophytes* genera discovered in the samples are clearly freshwater ones, and the depth of the watersurely didn't exceed 5 meters. Transportation of these forms is possible but improbable; the *Charophytes* are too numerous for all of them to be transported.

#### 5. Results

In this study, eight genera with fourteen species of planktic foraminifera and thirty genera with five species of benthic foraminifera were defined in the samples collected from the measured sections of the Celtek and Armutlu formations of the Corum-Cankırı basin of northern Turkey.

In addition, two planktic foraminiferal biozones and one benthic foraminiferal biozones were described from the Celtek and Armutlu Formations in the Amasya region. These are named the Lower Eocene *Morozovella aragonensis-Acarinina pentacamerata* and the Middle Eocene *Nummulites laevigatus* biozone. These zones were determined using the samples of the Armutlu Formation only.



Figure 8. Schematic depositional environments of Çeltek and Armutlu Formations

Concerning the charophytes, a precise datation remains hypothetical, because of the taxonomic uncertainties discussed before. According to obtained data, a Lower Ypresian age (Pivetaui to *Disermas* zone, see Riveline et al., 1996) can be attributed to the sediments if:

- the Nitellopsis is effectively an N. sigali one
- the small Charophytes belong to Peckichara
- the Harrisichara is an H. tiquetra one.

Paleontological and sedimentological data indicate that the Celtek and Armutlu formations were deposited in delta, lagoon and shallow marine environments.

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### 7. References

- Akkiraz, M. S., M. S. Kayseri & Akgun, F. (2008). Paleoecology of coal-bearing Eocene sediments in central Anatolia (Turkey) based on Quantitative Palynological Data. Turkish Journal of Earth Sciences, 17, 317-360.
- Atalay, Z. (2001). Amasya yoresindeki linyitli Celtek Formasyonu'nun stratigrafisi, fasiyes ve cokelme ortamı ozellikleri. Turkiye Jeoloji Bulteni, 44 (2), 1-22.
- Beckmann, J. P., I. El-Heiny, M. Kerdany, R. Said & Viotti, C. (1969). Standart planktic zones Egypt. Proceed First Inter. Conf. Planktonic mikrofossils, 1, 92-103.
- Berggren, W. A., D. V. Kent, C. C. Swisher & Aubry, M. P. (1995). A revised Cenozoic geochronology and chronostratigraphy. Society Econom. Paleont. and Mineral., 54, 129-212.
- Blumenthal, M. M. (1937). Amasya Vilayetine tabi Celtek'teki Linyitli Arazinin Jeolojik Tesekkulati Hakkinda rapor. Maden Tetkik ve Arama Enstitusu, 157, Ankara.
- Bolli, H. M. (1957a). The genera Globigerina and Globorotalia in the Paleocene-Lower Eocene Lizard Springs Formation of Trinidad. B. W. I. Bull. U. S. Natl. Mus., 215, 61-81.
- Bolli, H. M. (1957b). Planktonic foraminifera from the Eocene Navet and San Fernando formations of Trinidad. B. W. I. Bull. U. S. Natl. Mus., 215, 72-155.
- Bolli, H. M. (1966). Zonation of Cretaceous to Pliocene marine sediments based on planktonic foraminifera. Boletin Informativo Associacion Venezolana de Geologia, Mineriay Petroleo, 9/1, 3-32.
- Ellis, B. F. & Messina, A. R. (1966). Catalogue of index foraminifera. Special Publication 2, The American Museum of Natural History, New York.
- Ejel, F. (1967). Zones stratigraphiques du Paléogéne et probléme de la limited Eocéne supériuer dansla région de Damas (Syrie). Proceed First Intern. Cof. Planktonic microfossils, 2, 175-181.
- Eris, E. (1996). Eosen yasli Celtek (Amasya) komurlerinin kimyasal-petrografik ozellikleri, oluşum ortamı ve ekonomik potansiyelinin incelenmesi. Yuk. Lis. Tezi, H.U., Fen. Bil. Enst., 63 s.
- Gumussu, M. (1978). Amasya ili, Merzifon ve Suluova İlceleri komur jeolojisi. MTA Derleme Rapor, 7063, Ankara.
- Hazerfan, C. (1974). (Amasya-Suluova) Celtek komur isletmesi civarinin jeolojik raporu. MTA Derleme Rapor, 6137, Ankara.
- Karayigit, A. İ., E. Eris & Cicioglu, E. (1996). Coal geology, chemical and petrographical characteristics and implications for coal bed methane development of subbitumious coals from the Sorgun and Suluova basins, Turkey. In: Methane and Coal Geology, Geol. Soc. Spec. Publ., 97, 327-340.
- Kayseri, M. S. & Akgun, F. (2003). Palynofloristic correlation of Neogene sediments in western and Central Anatolia (TURKEY). Neogene Climate Evolution in Eurasia Annual Meeting (NECLIME), 12–13.
- Koc, C. (2002). Suluova (Amasya) kuzeybatisindaki Eosen cokellerinde sedimantolojik incelemeler. Yuksek Lisans Tezi, Firat Universitesi Fen Bilimleri Enstitusu, 52 s.
- Koc, C. & Turkmen, I. (2002). Suluova (Amasya) kuzeyindeki komurlu Eosen cokellerinin sedimantolojik ozellikleri. Hacettepe Universitesi, Yerbilimleri Bulteni, 26, 101-117.
- Koc, C., I. Turkmen & Kaya, M. (2002). Celtek ve Armutlu formasyonlarinin stratigrafik ozellikleri, Suluova (Amasya) Kuzeyi. Geosound/Yerbilimleri, 40/41, 81-96.
- Krasheninnikov, V. A. (1965). Zonal Stratigraphy of the Paleogene in the eastern Mediterranean. Acad. Nauk. SSSR Geol. Inst., 133, 1-76.
- Olsson, R. K., C. Hemleben, W. A. Berggren & Huber, B. T. (1999). Atlas of Paleocene Planktonic foraminifera. Smithsonian Contributions to Paleobiol., 85, 1-252.

Ozdemir, I. & Pekmezci, F. (1983). Amasya-Suluova-Celtek Linyit Sahalari Jeolojisi ve Sondaj Aramalari Raporu. Maden Tetkik Arama Enstitusu, Atom Komur Dairesi, 428 s, Ankara.

Postuma, J. A. (1971). Manual of planktonic foraminifera. Elsevier Publishing Company, Amsterdam, 420 pp.

- Riveline, J., J. P. Berger, M. Feist, C. Martin-Closas, M. Schudack & Soulie-Märsche, I. (1996). European Mesozoic-Cenozoic charophyte biozonation. Bull. Soc. Géol. France, 167/3, 453-468.
- Sari, I. (2008). Armutlu (Suluova-Amasya) linyitlerinin bazi ozellikleri, rezerv hesabi, ve ekonomik onemi. Yuk. Lis. Tezi, Selcuk Univ. Fen Bil. Enst., 83 s.
- Sengor, A. M. C. & Yilmaz, Y. (1983). Turkiye'de Tetis'in evrimi: Levha tektonigi acisindan bir yaklasim. Turkiye Jeoloji Kurumu Yerbilimleri Ozel Dizisi, 1, 75 s.
- Serra-Kiel, J., L. Hottinger, E. Caus, K. Drobne, C. Ferrandez, A. K. Jauhri, G. Less, R. Pavlovec, J. Pignatti, J. M. Samso, H. Schaub, E. Sirel, A. Strougo, Y. Tambareau, J. Tosquella & Zakrevskaya, E. (1999). Larger foraminiferal biostratigraphy of the Tethyan Paleocene and Eocene. Bulletin Sociéty Géology, France, Paris, 169, 281-299.
- Toumarkine, M. (1978). Planktonic foraminiferal biostratigraphy of the Paleogene of sites 360 to 364 and the Neogene of sites 362a, 363 and 364 Leg 40. Initial Rep. Deep Sea Drilling Project, 40, 679-721.
- Toumarkine, M. & Luterbacher, H. P. (1985). Luterbacher Paleocene and Eocene planktic foraminifera. In: Bolli H.M, Saunders J. B, Perc-Nielsen K. (Eds). Plankton Stratigraphy, Cambridge University Press, Cambridge, 87-154.

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## Plate I

*Acarinina pentacamerata* (SUBBOTINA) Figure 1. Umbilical side, sample no. 4, New Celtek section.

*Acarinina soldadoensis soldadoensis* (BRONNIMANN) Figure 2. Vertical section, X100, sample no. 1, New Celtek section.

*Morozovella acuta* (TOULMIN) Figure 3. Vertical section, X100, sample no. 8, Kalayli Hill section.

*Morozovella aragonensis* (NUTTALL) Figure 4. Spiral side, sample no. 1, New Celtek section.

Figure 5. Side view, sample no. 1, New Celtek section.

## Morozovella subbotinae (MOROZOVA)

Figure 6. Vertical section, X100, sample no. 11, Armutlu Village section.



PLATE I

#### Plate II

#### Nummulites laevigatus (BRUGUIERE)

Figure 1. Equatorial section, microspheric form, X3, sample no. 5, Sirikli Hill section.

Figure 2. Axial section, microspheric form, X2.5, sample no. 5, Sirikli Hill section.

Figure 3. Surface view, microspheric form, X3, sample no. 5, Sirikli Hill section.

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Figure 5. Axial section, macrospheric form, X7, sample no. 5, Sirikli Hill section.

#### Nummulites lehneri SCHAUB

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Figure 11. Equatorial section, macrospheric form, X6.5, sample no. 4, Sirikli Hill section.

#### Nummulites uranensis(dé la HARPE)

Figure 12. Equatorial section, microspheric form, X6.5, sample no. 5, Sirikli Hill section.

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Figure 15. Equatorial section, macrospheric form, X8, sample no. 5, Sirikli Hill section.

Figure 16. Axial section, macrospheric form, X9, sample no. 5, Sirikli Hill section.



PLATE II